

數值分析於大地工程上的應用
Application of numerical simulation in
geotechnical engineering

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魏敏樺 博士
國立台北科技大學
土木工程系
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Outlines:

1. Introduction
2. Theoretical background
3. Verification of formulation
4. Real case study
5. Conclusions

Introduction

2008/9/28, 7.22pm

Area: 20~40 m (W) x 220 m (L)

Depth: 4~7 m (Top-soil & weathered rock)

Volume: 30,000 m³

Tower T16



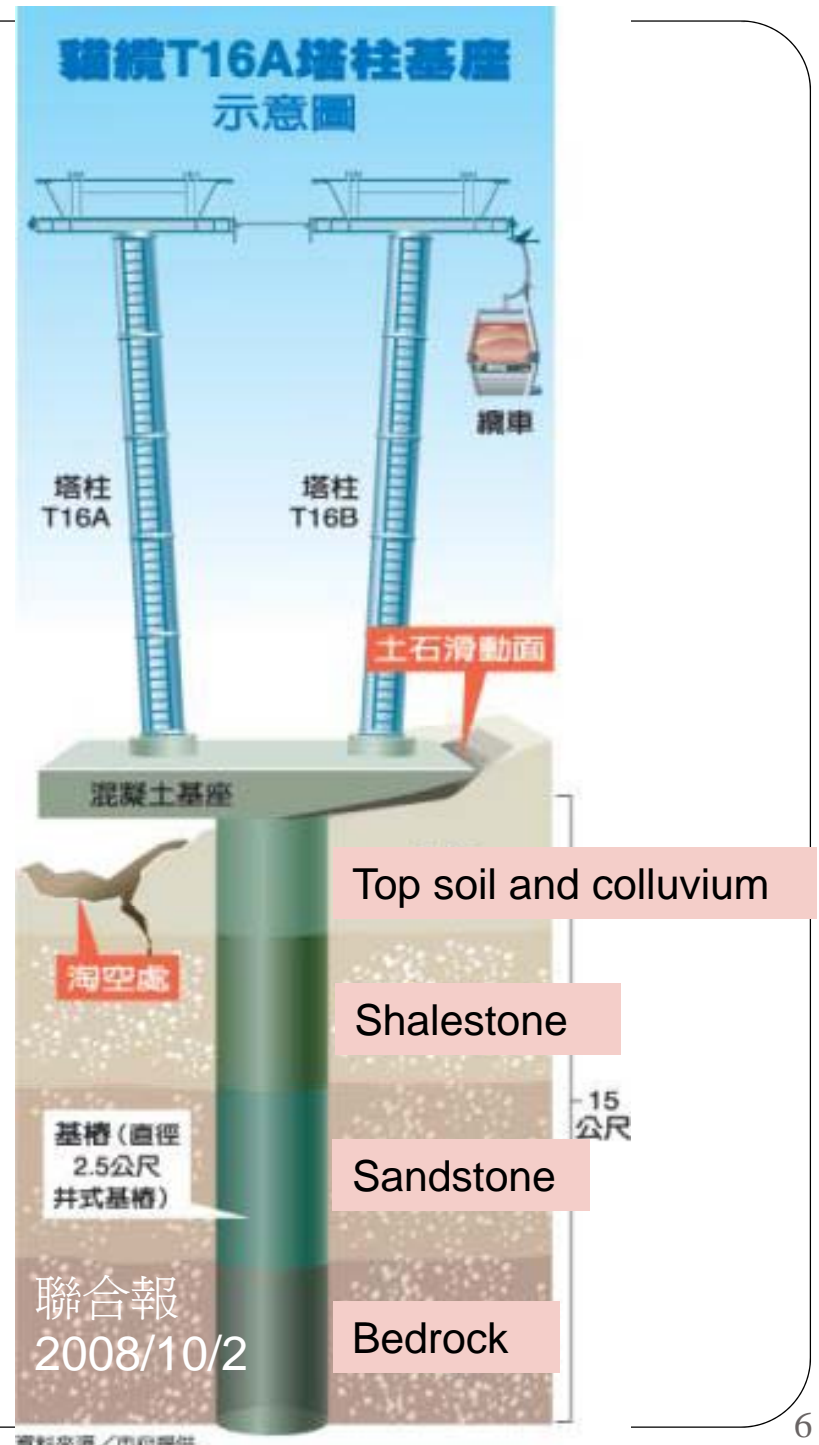
政大御花園後方

Maokong-Gondola (cable car) Tower-T16 downslope

資料來源：大地工程學會（2009），「木柵地區-大地工程與地質野外考察手冊」。
<<http://www.tgs.org.tw/tgs/index.php/-mainmenu-33/80--->>

齊柏林





Theoretical Background

Equilibrium eqn for saturated soil:

$$\begin{aligned} \frac{\partial \sigma_x}{\partial x} + \frac{\partial \tau_{yx}}{\partial y} + \frac{\partial \tau_{zx}}{\partial z} &= 0 \\ \frac{\partial \tau_{xy}}{\partial x} + \frac{\partial \sigma_y}{\partial y} + \frac{\partial \tau_{zy}}{\partial z} &= 0 \\ \frac{\partial \tau_{xz}}{\partial x} + \frac{\partial \tau_{yz}}{\partial y} + \frac{\partial \sigma_z}{\partial z} &= -\gamma \end{aligned}$$

Effective stress for unsaturated soil:

$$\sigma' = (\sigma - u_a) + \beta(u_a - u_w)$$

Soil structure constitutive relations for unsaturated soil:

$$\begin{aligned} \sigma_x &= \frac{E}{(1 + \mu)} \varepsilon_x + \frac{E\mu\varepsilon_v}{(1 + \mu)(1 - 2\mu)} + u_a - \beta(u_a - u_w) \\ \sigma_y &= \frac{E}{(1 + \mu)} \varepsilon_y + \frac{E\mu\varepsilon_v}{(1 + \mu)(1 - 2\mu)} + u_a - \beta(u_a - u_w) \\ \sigma_z &= \frac{E}{(1 + \mu)} \varepsilon_z + \frac{E\mu\varepsilon_v}{(1 + \mu)(1 - 2\mu)} + u_a - \beta(u_a - u_w) \end{aligned}$$

Compatibility equation:

$$\left\{ \begin{aligned} \varepsilon_x &= \frac{\partial u}{\partial x} ; \varepsilon_y = \frac{\partial v}{\partial y} ; \varepsilon_z = \frac{\partial w}{\partial z} \\ \gamma_{yz} &= \left(\frac{\partial u}{\partial z} + \frac{\partial w}{\partial z} \right) ; \gamma_{xz} = \left(\frac{\partial u}{\partial z} + \frac{\partial w}{\partial x} \right) \\ \gamma_{xy} &= \left(\frac{\partial v}{\partial x} + \frac{\partial u}{\partial y} \right) ; \tau = G\gamma \end{aligned} \right.$$

$$G = \frac{E}{2(1+\mu)}; \beta = \frac{E}{H(1-2\mu)}$$

Equilibrium DE in terms of $(u_a - u_w)$ and ε_v for unsaturated soil:

$$G\nabla^2 u + \frac{G}{1-2\mu} \frac{\partial \varepsilon_v}{\partial x} - \beta \frac{\partial (u_a - u_w)}{\partial x} + \frac{\partial u_a}{\partial x} = 0$$

$$G\nabla^2 v + \frac{G}{1-2\mu} \frac{\partial \varepsilon_v}{\partial y} - \beta \frac{\partial (u_a - u_w)}{\partial y} + \frac{\partial u_a}{\partial y} = 0$$

$$G\nabla^2 w + \frac{G}{1-2\mu} \frac{\partial \varepsilon_v}{\partial z} - \beta \frac{\partial (u_a - u_w)}{\partial z} + \frac{\partial u_a}{\partial z} = -\gamma$$

$$\nabla^2 = \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2} + \frac{\partial^2}{\partial z^2}$$

$u, v, w, (u_a - u_w)$ = four unknowns to solve for but with only three equations

Continuity equation eqn for water flow through an unsaturated soil:

$$\nabla \cdot \left[k_w^x \frac{\partial h}{\partial x} + k_w^y \frac{\partial h}{\partial y} + k_w^z \frac{\partial h}{\partial z} \right] = \left(\frac{m_1^w}{m_2^s} \right) \frac{\partial \varepsilon_v}{\partial t} + \gamma_w \left(m_2^w - \frac{m_1^w m_2^s}{m_1^s} \right) \frac{\partial h}{\partial t}$$

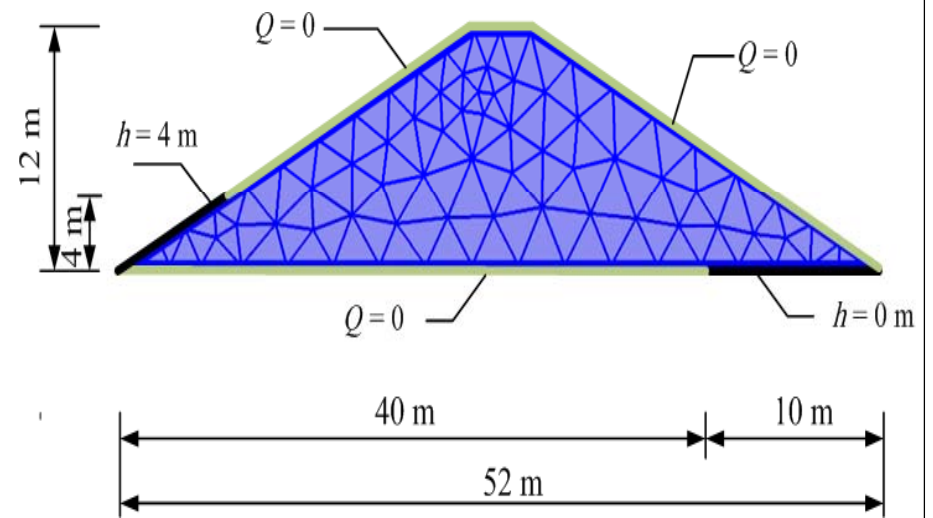
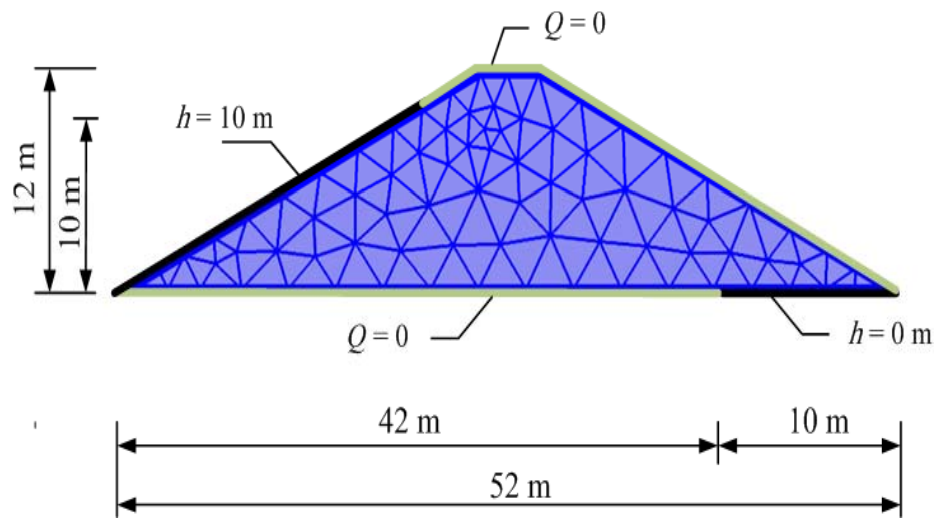
$$m_1^s = \frac{d\varepsilon_v}{d(\sigma_{\text{mean}} - u_a)} = \frac{1}{(1 + e_0)} \frac{de}{d(\sigma_{\text{mean}} - u_a)}$$

$$m_2^s = \frac{d\varepsilon_v}{d(u_a - u_w)} = \frac{1}{(1 + e_0)} \frac{de}{d(u_a - u_w)}$$

$$m_1^w = \frac{\partial \theta_w}{\partial(\sigma_{\text{mean}} - u_a)} = \frac{S}{(1 + e_0)} \frac{e}{\frac{dS}{de} d(\sigma_{\text{mean}} - u_a)}$$

$$m_2^w = \frac{\partial \theta_w}{\partial(u_a - u_w)} = \frac{S}{(1 + e_0)} \frac{e}{d(u_a - u_w)} + \frac{e}{(1 + e_0)} \frac{dS}{d(u_a - u_w)}$$

Verification



$$\frac{\partial}{\partial x} \left(k_x \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left(k_y \frac{\partial h}{\partial y} \right) = m_2^w \rho_w g \frac{\partial h}{\partial t}$$

van Genuchten (1980)

回歸係數

a

n

m

儲水方程式：

$$m_2^w = -\frac{\theta_s}{(1 + (a\psi)^n)^m} (a\psi)^n \frac{mn}{\psi(1 + (a\psi)^n)}$$

4.391×10⁻²

46.754

0.036

$$\theta_s = 0.381$$

非飽和滲透係數方程式：

$$k_w = k_{sat} \left[\frac{[1 - (a\psi)^{n-1} [1 + (a\psi)^n]^{-m}]^2}{[1 + (a\psi)^n]^{\frac{m}{2}}} \right]$$

4.049×10⁻²

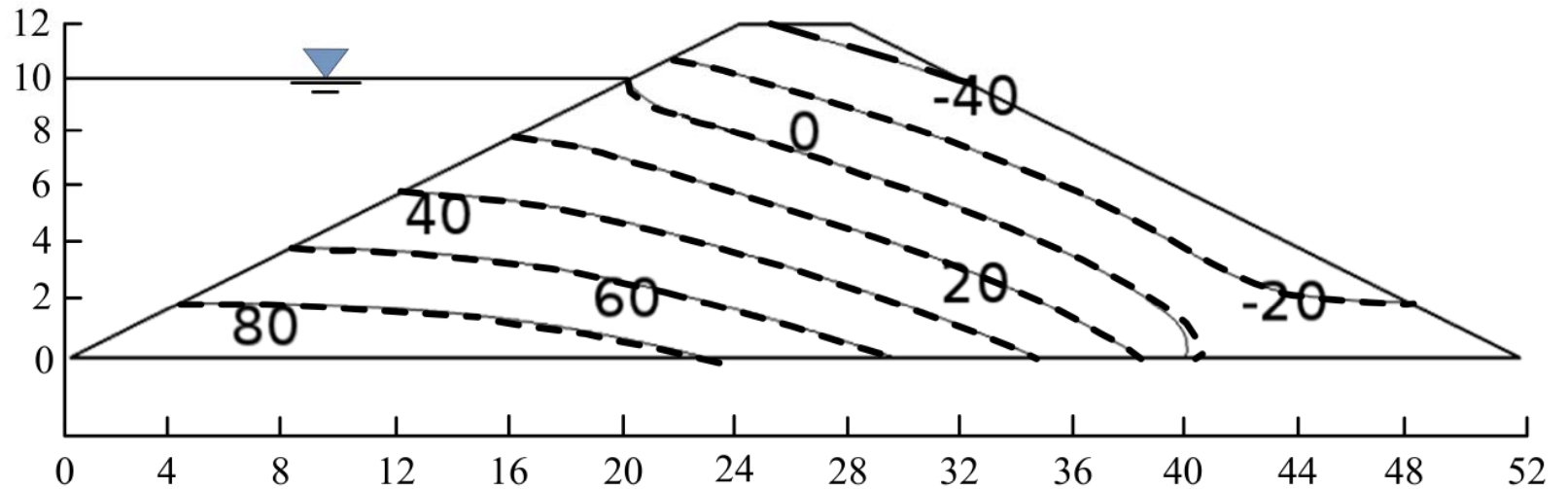
8.852

0.887

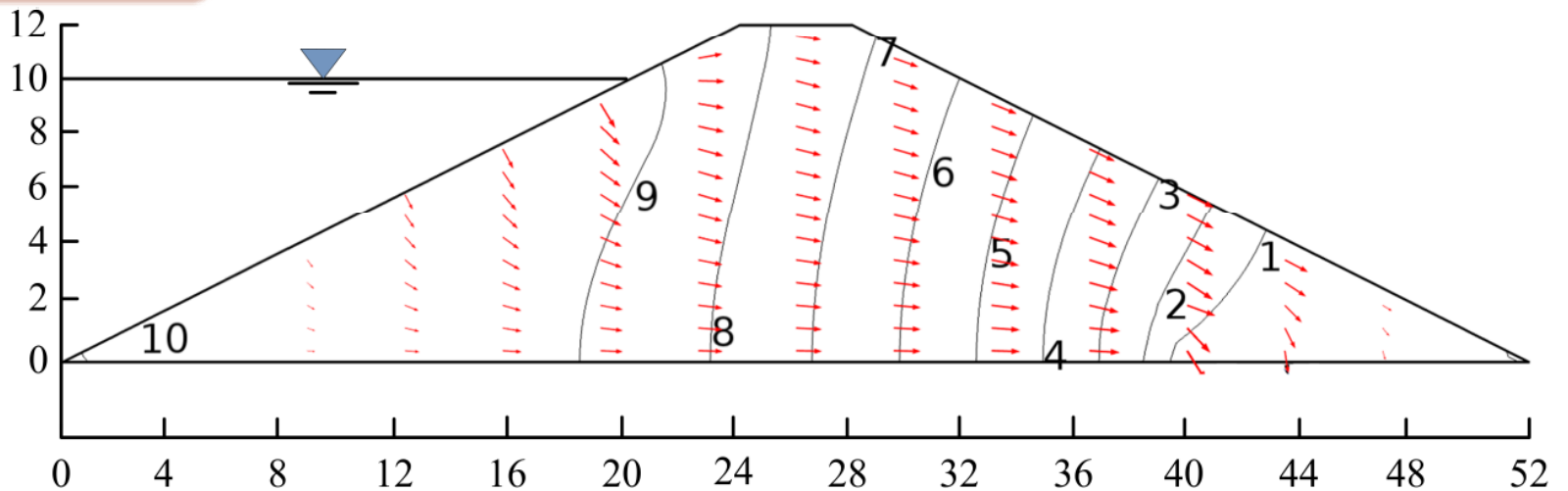
$$k_{sat} = 2.5 \times 10^{-7} \text{ (m/sec)}$$

穩態-蓄水面 10 m
(孔隙水壓力)

Thieu (2001) 虛線
Simulation 實線

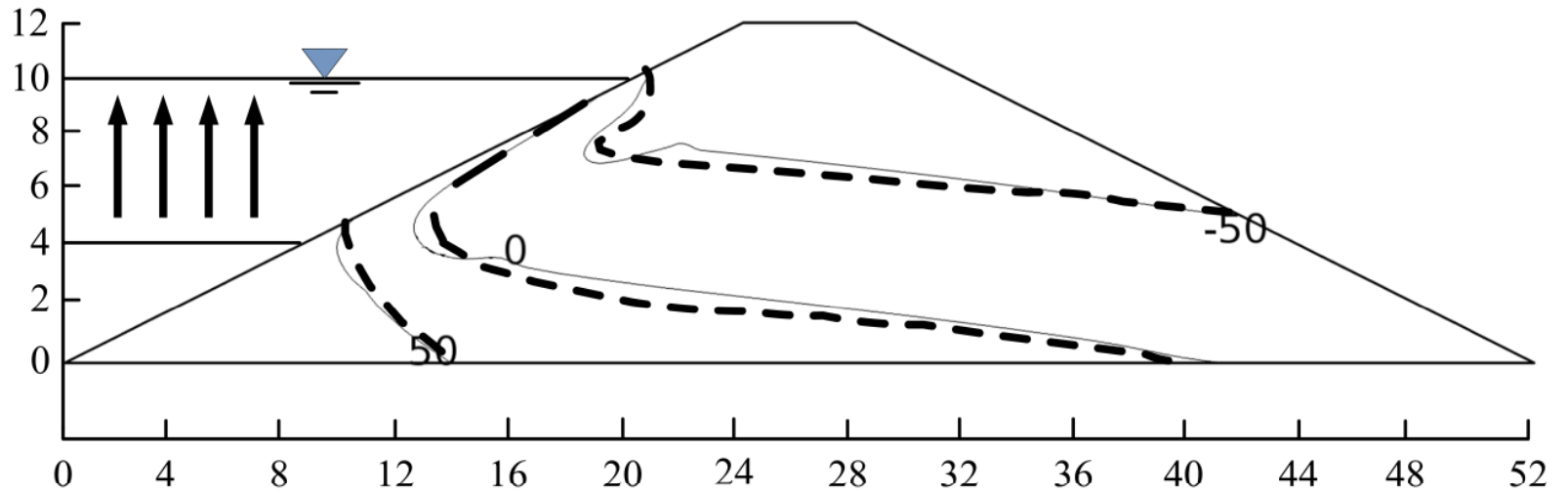


等勢能線(實線)
流線(箭頭)

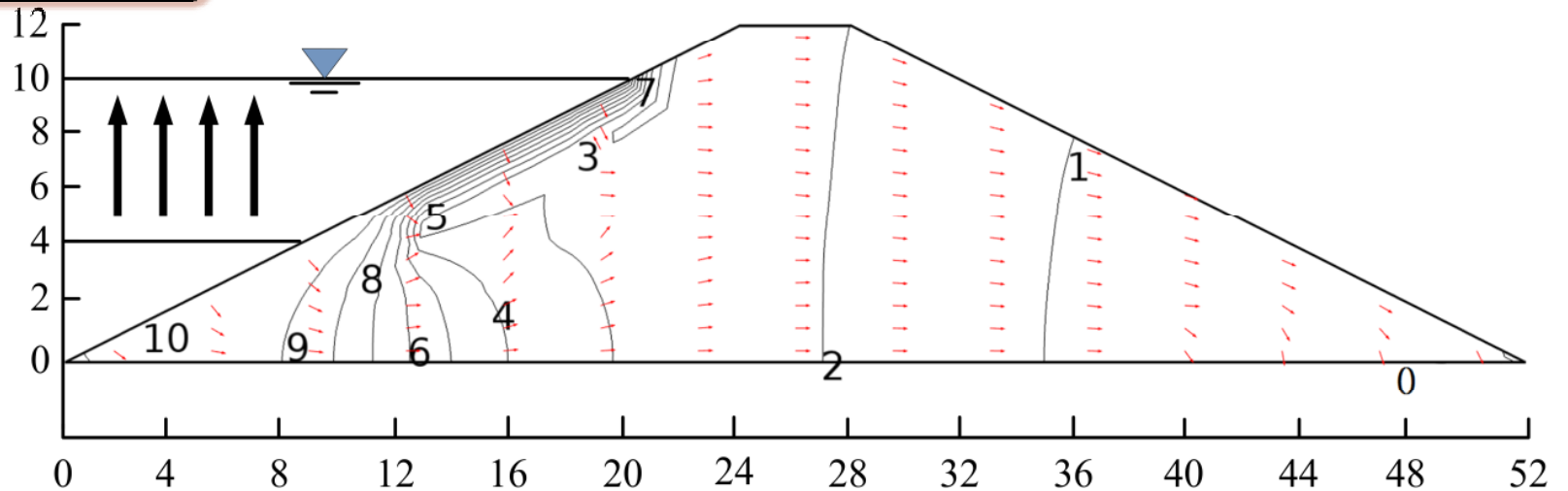


暫態-蓄水面 4 m 上升至 10 m
(孔隙水壓力)

Thieu (2001) 虛線
Simulation 實線

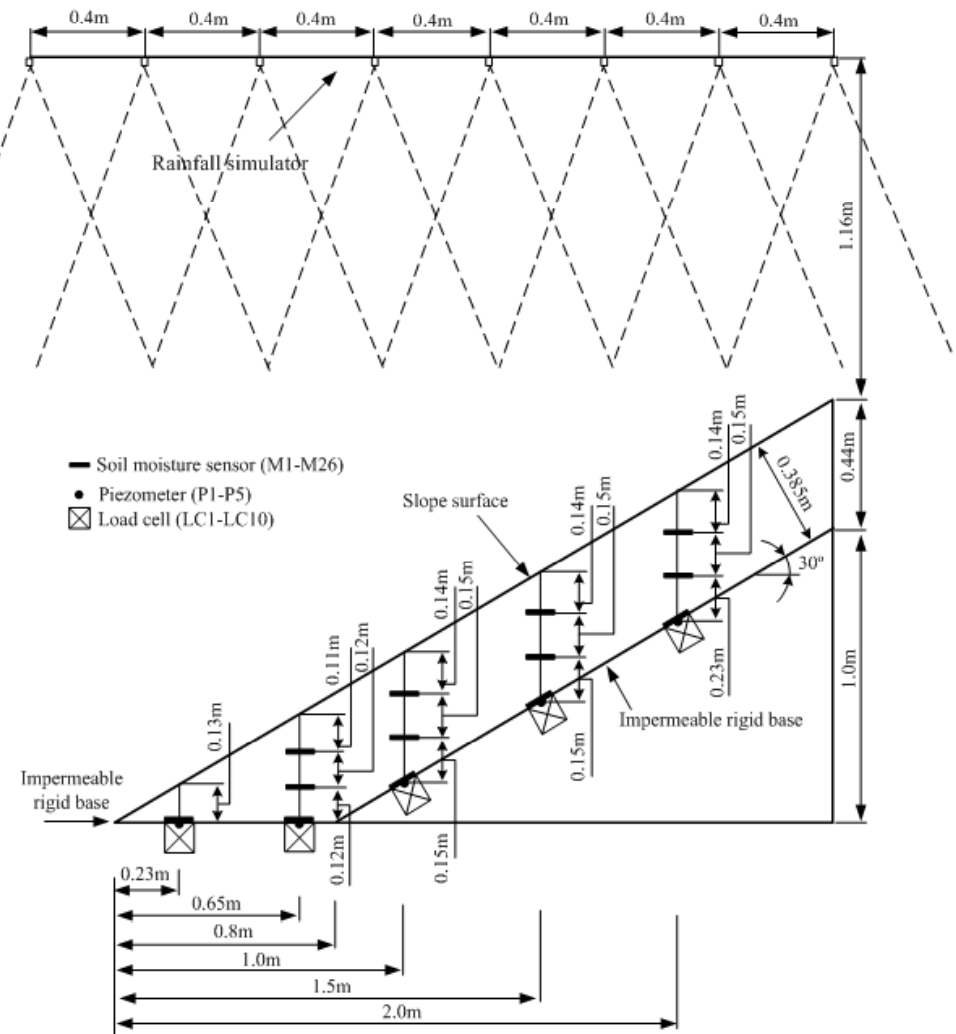


等勢能線(實線)
流線(箭頭)

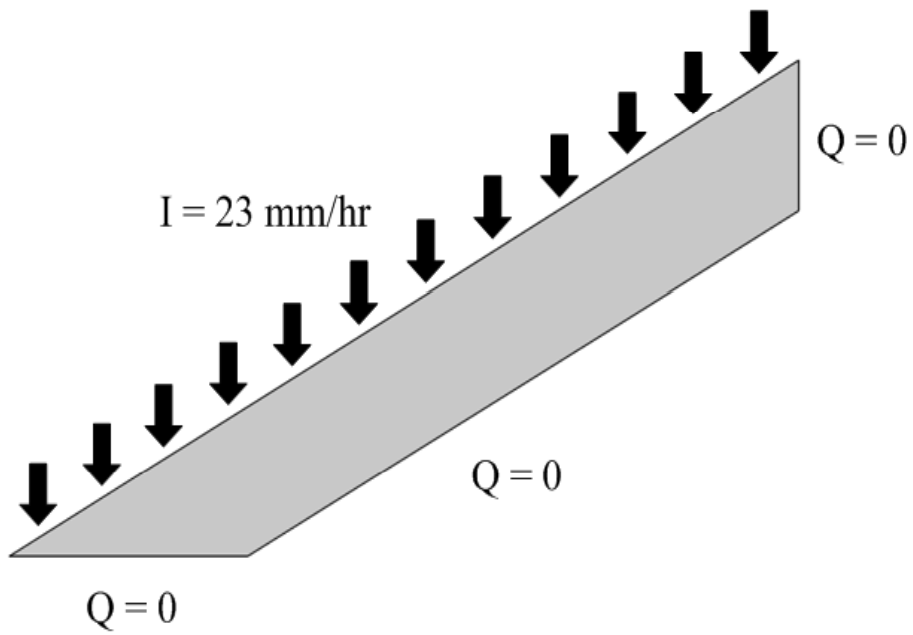


Simulation of Laboratory Model Test

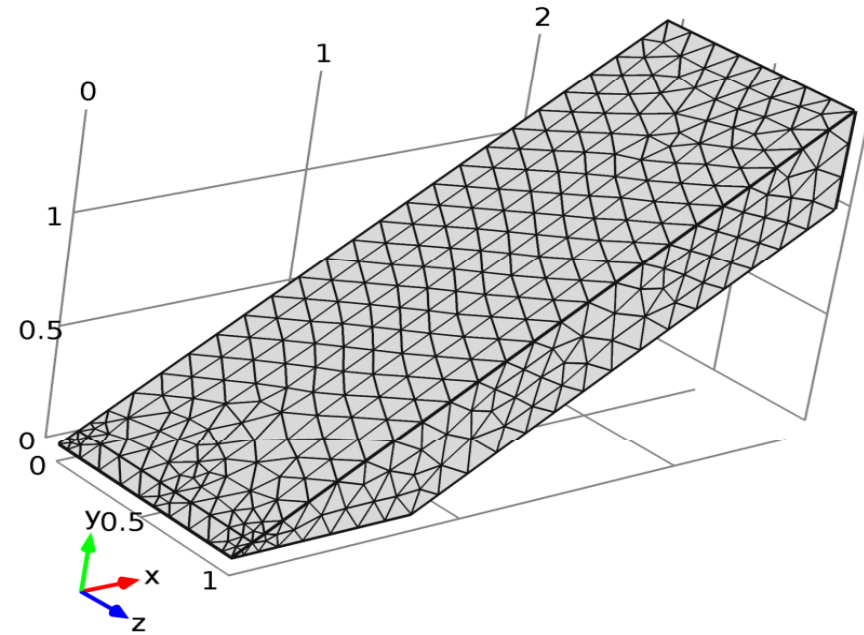
Simulate Lan's (2008) man-made rainfall slope model test and compared the simulated and experimental distribution of degree of saturation (water content).



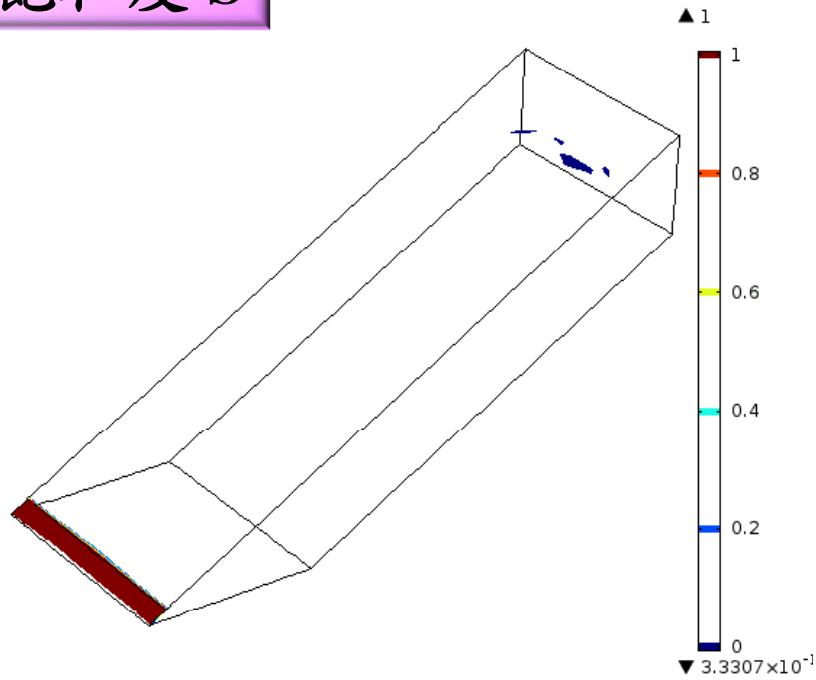
邊坡模型邊界狀態



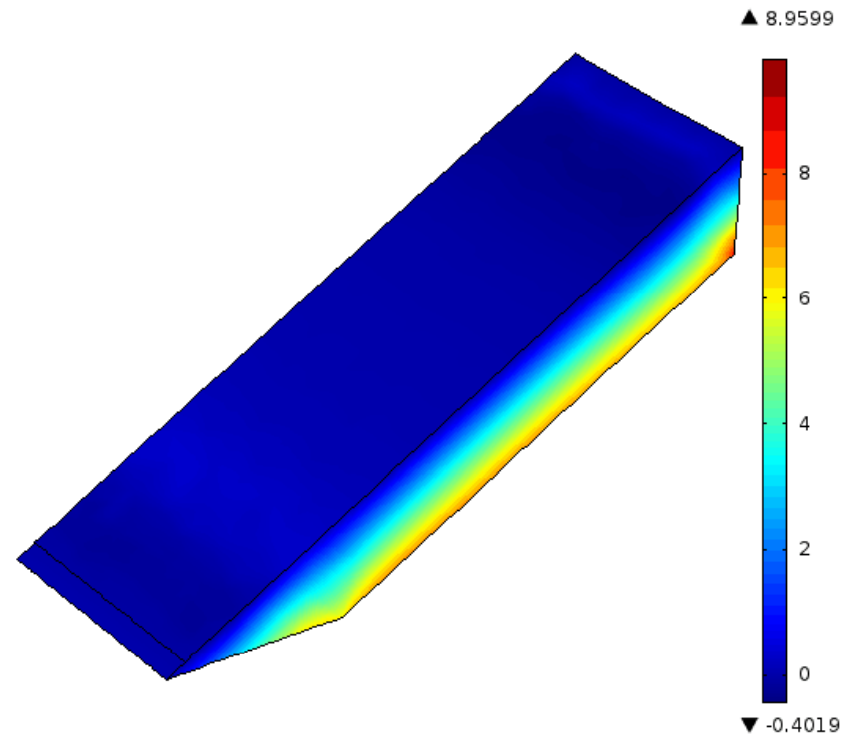
網格



飽和度 S

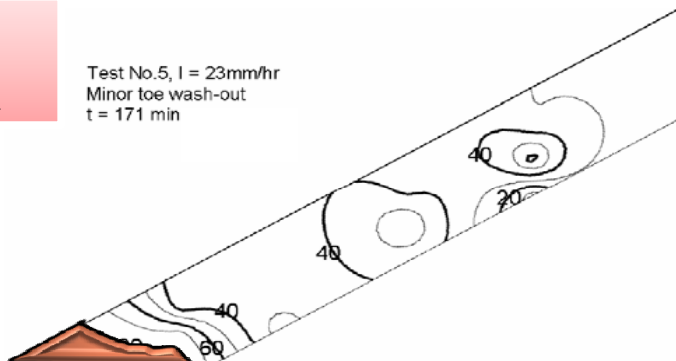


正向應力 kPa + 總位移

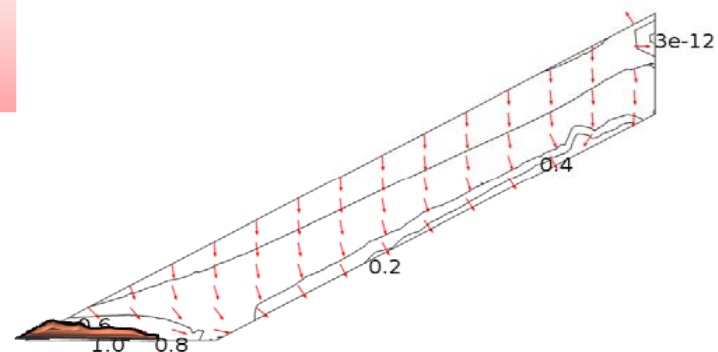


藍少村
171 min

Test No.5, $I = 23\text{mm/hr}$
Minor toe wash-out
 $t = 171\text{ min}$

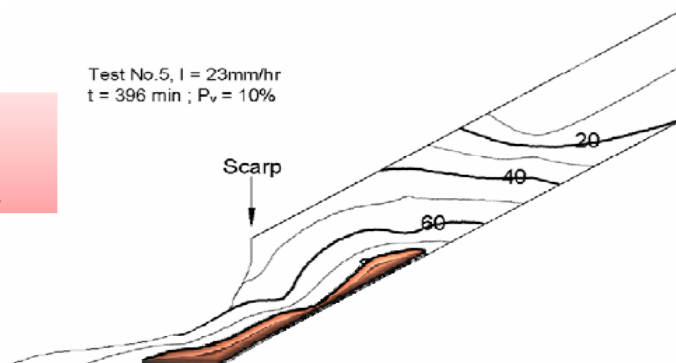


Simulation
171 min

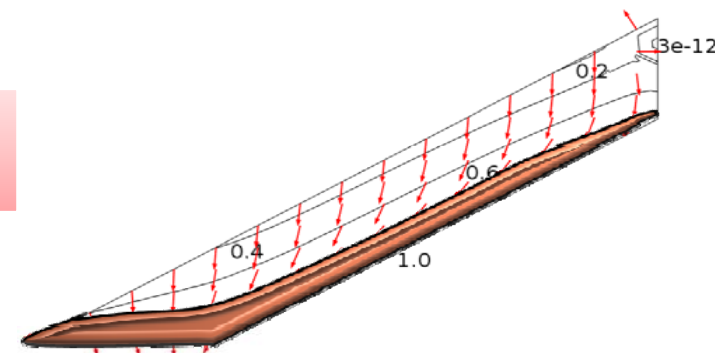


藍少村
396 min

Test No.5, $I = 23\text{mm/hr}$
 $t = 396\text{ min}$; $P_v = 10\%$

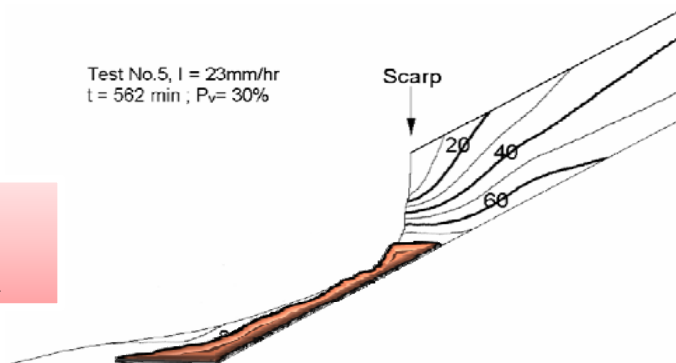


Simulation
396 min

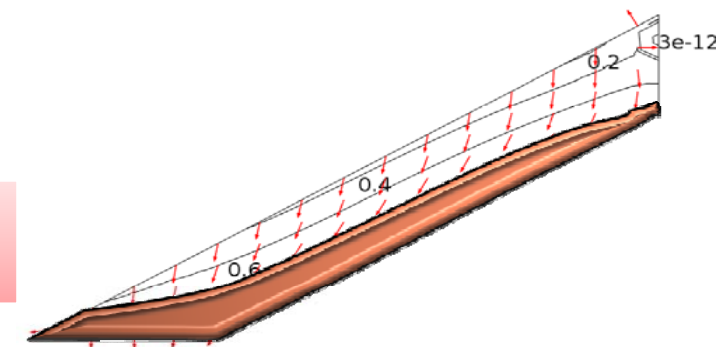


藍少村
562 min

Test No.5, $I = 23\text{mm/hr}$
 $t = 562\text{ min}$; $P_v = 30\%$

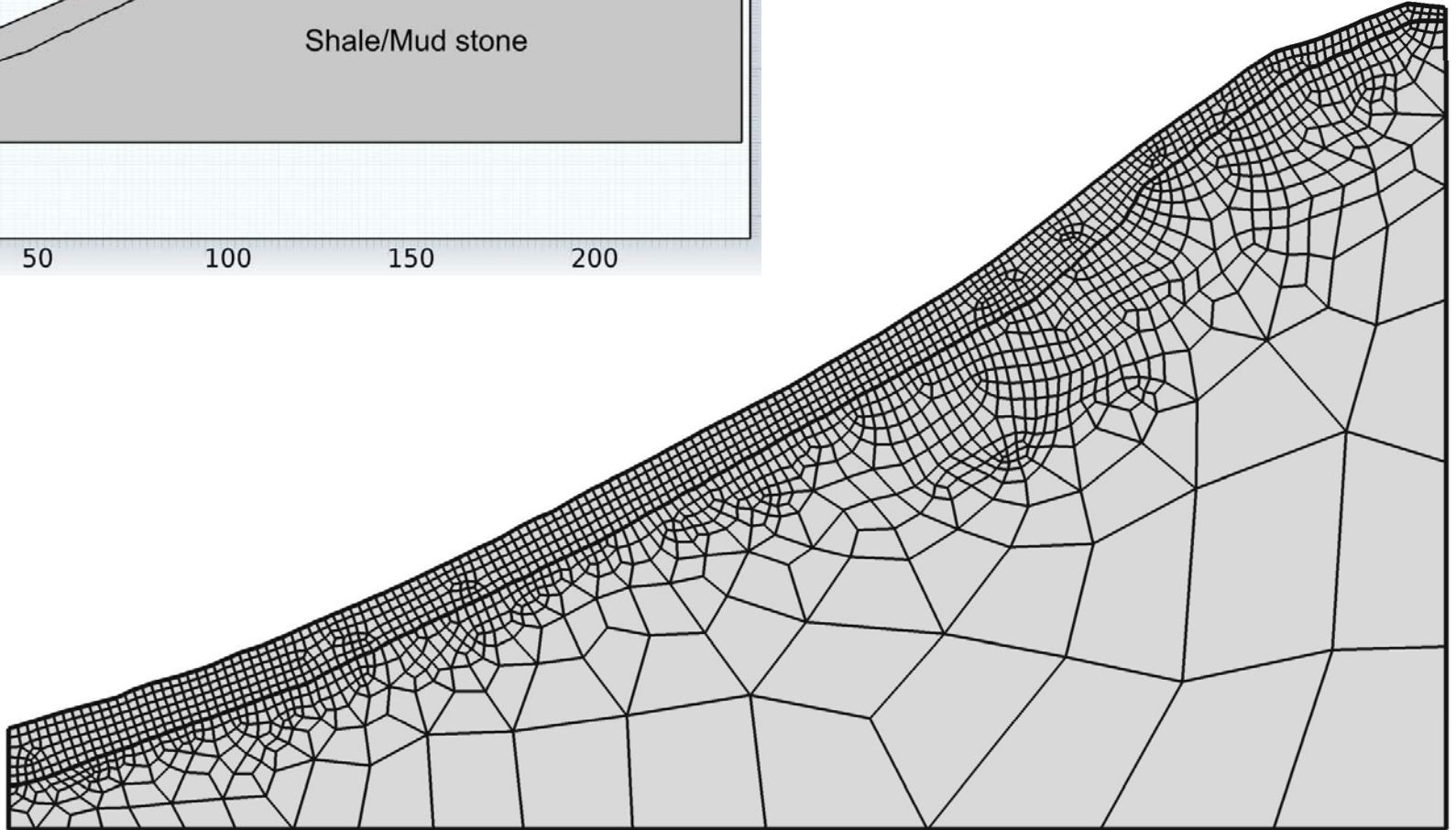
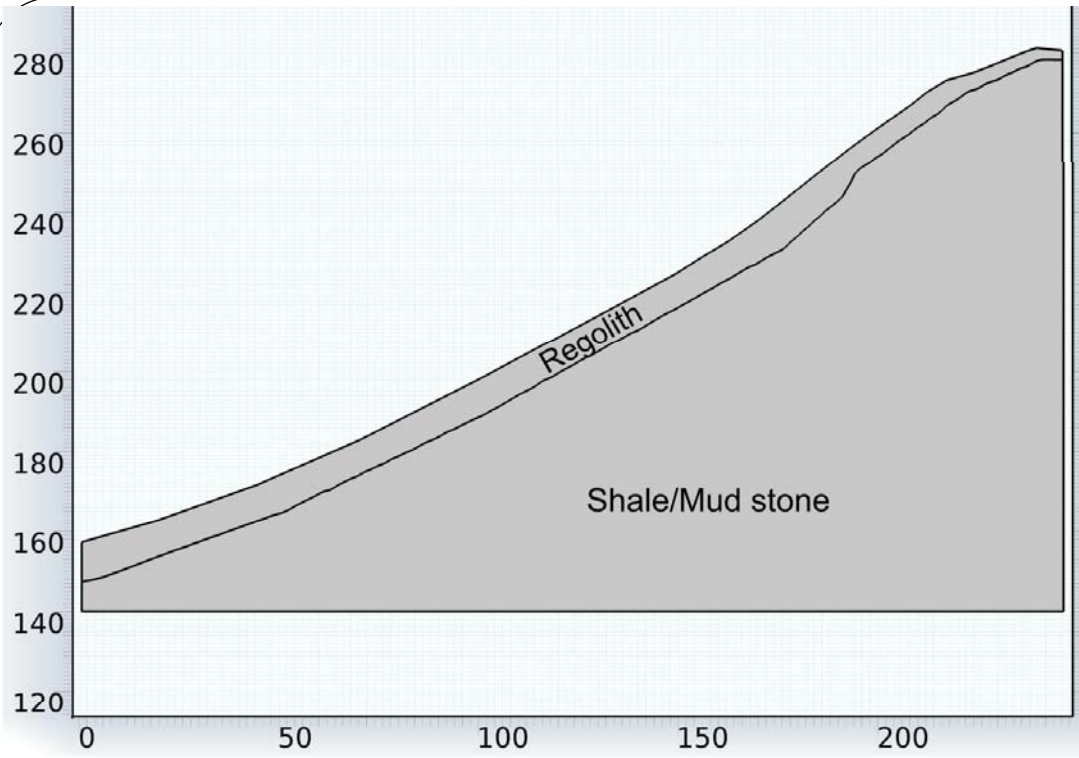


Simulation
562 min

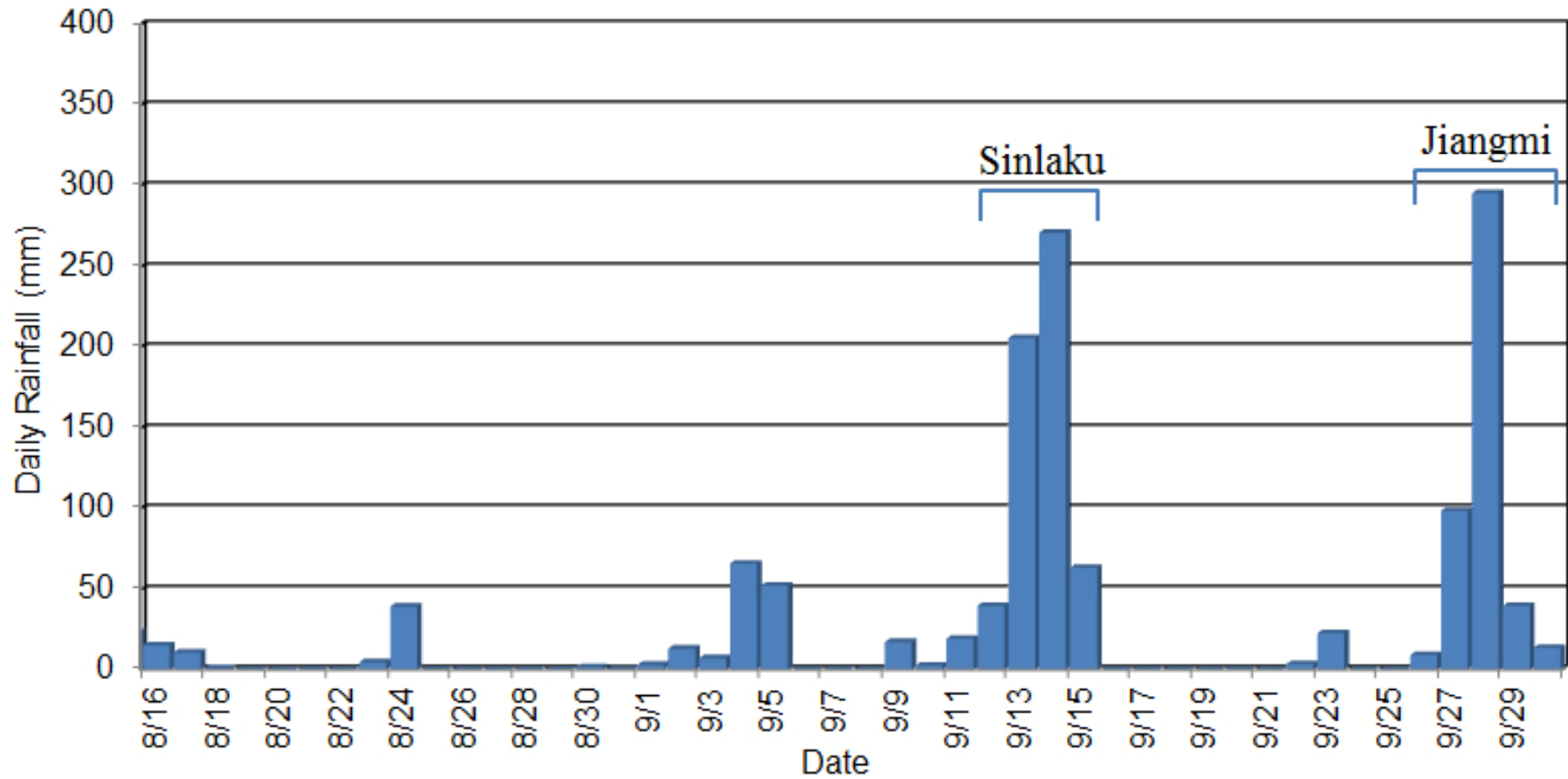


Real Case Study

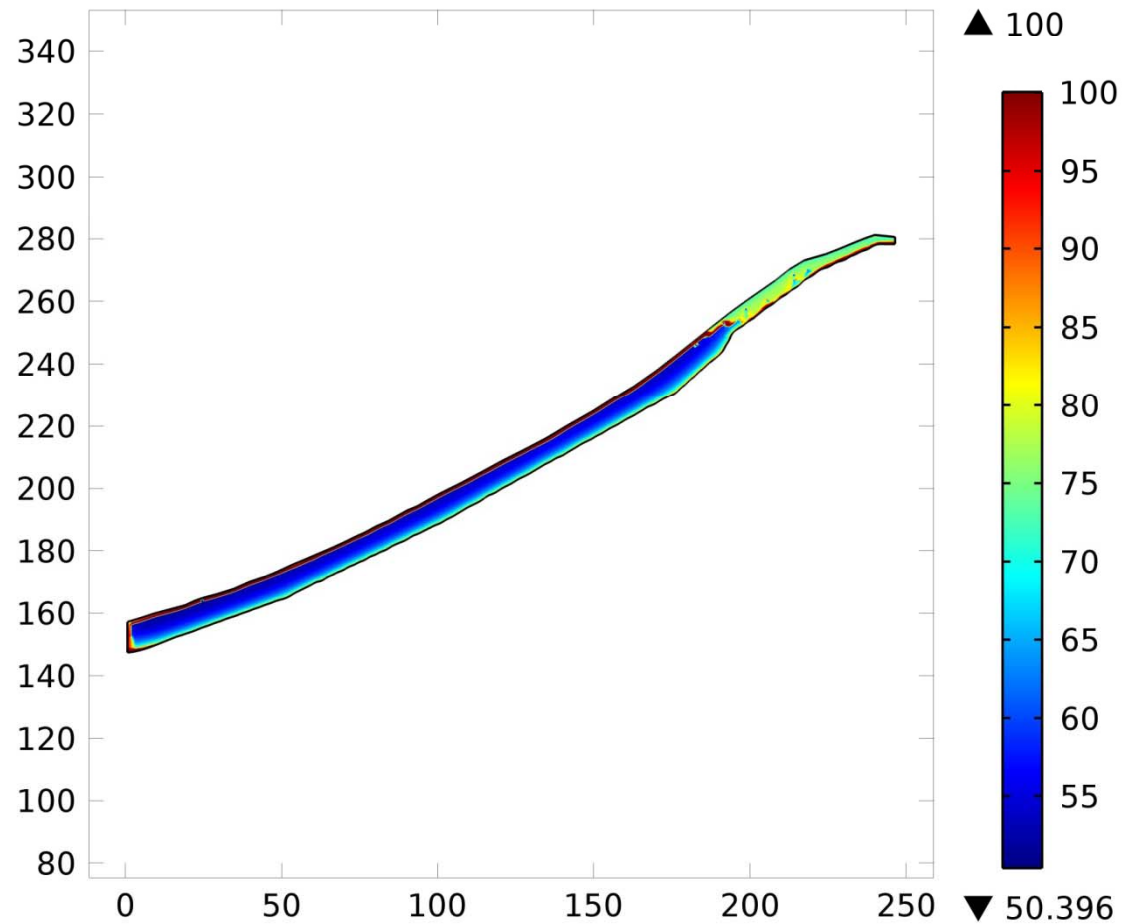
Simulation of Maokong Landslide



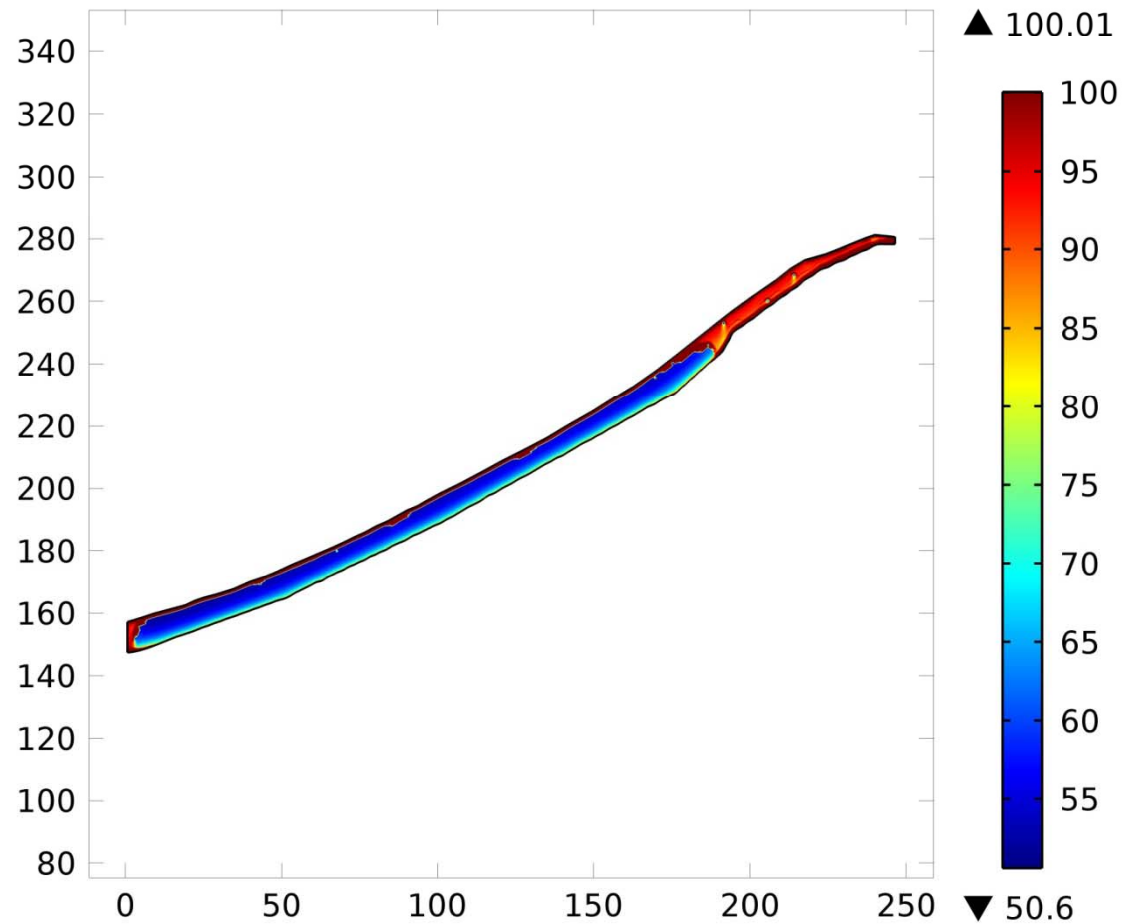
Daily rainfall recorded between 8/1-9/30



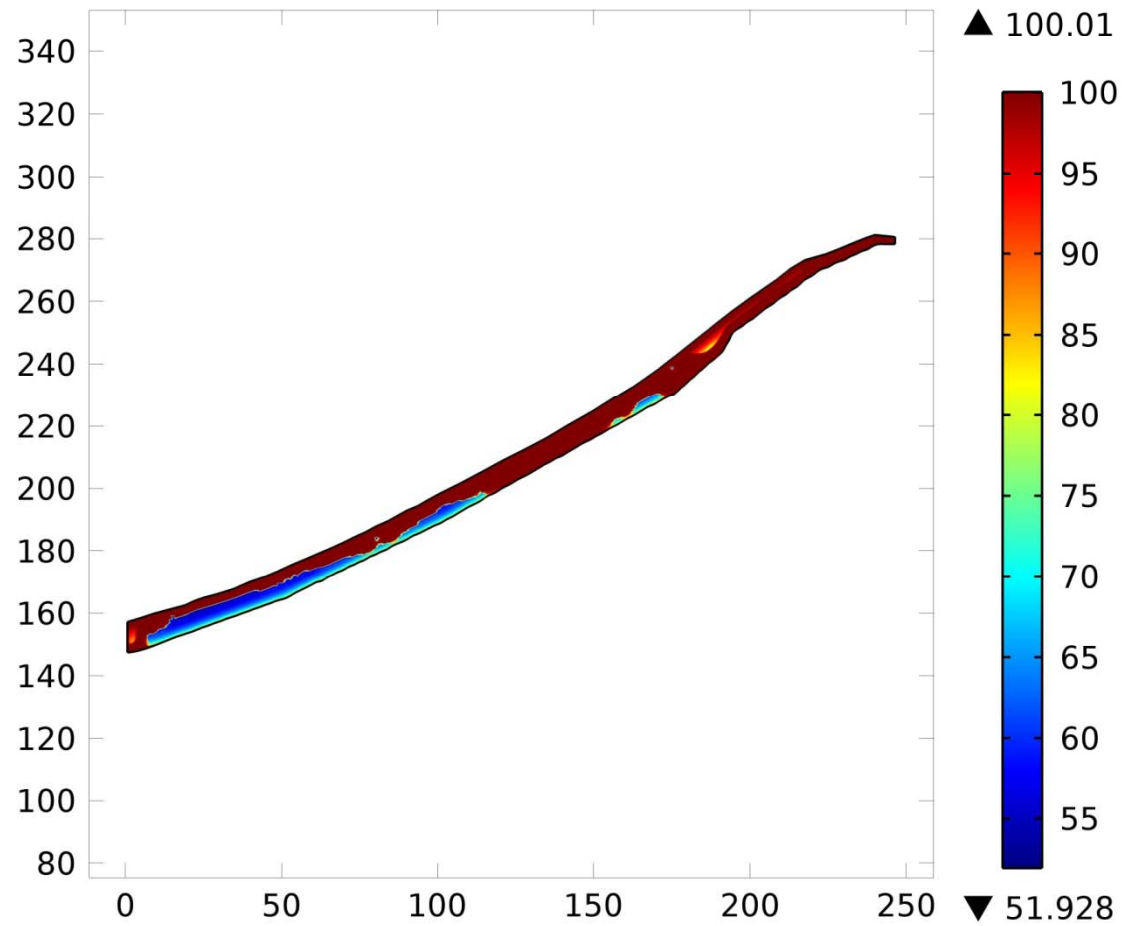
Degree of saturation on 8/30



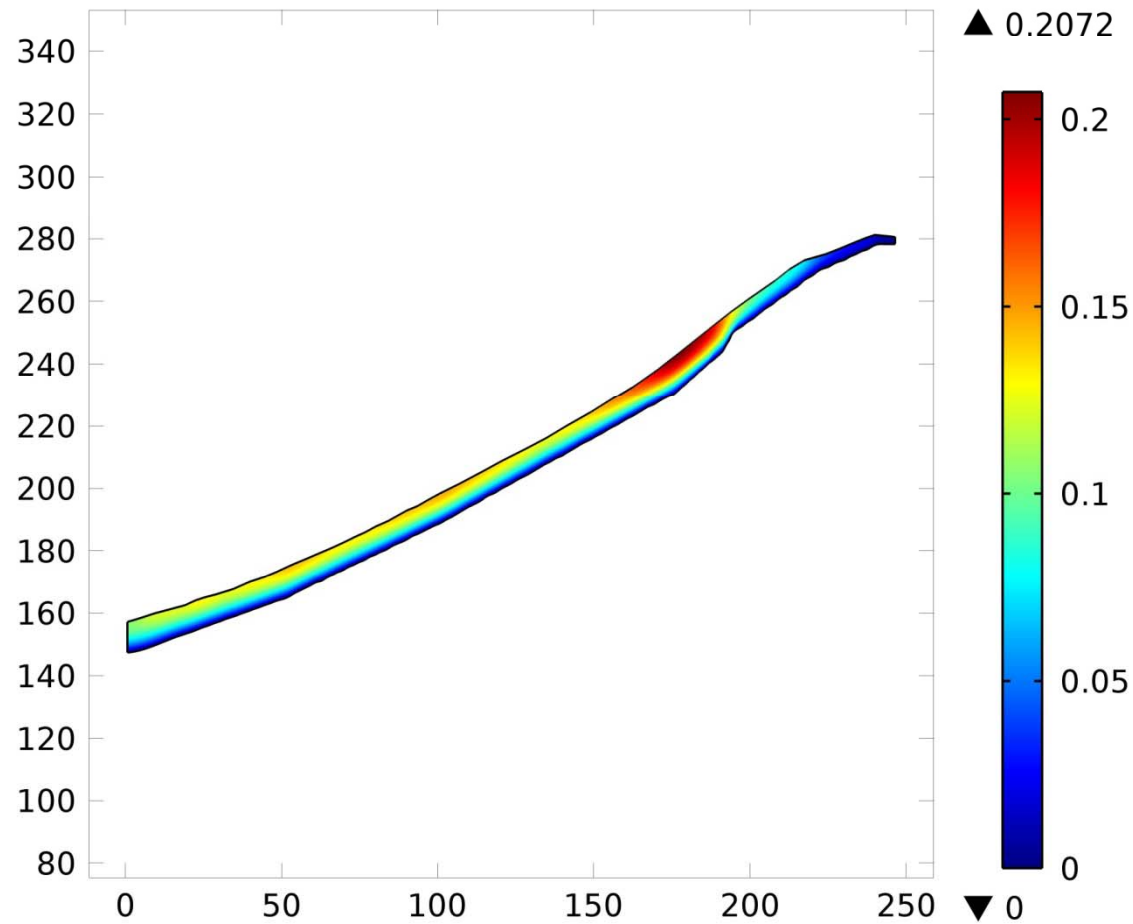
Degree of saturation on 9/14



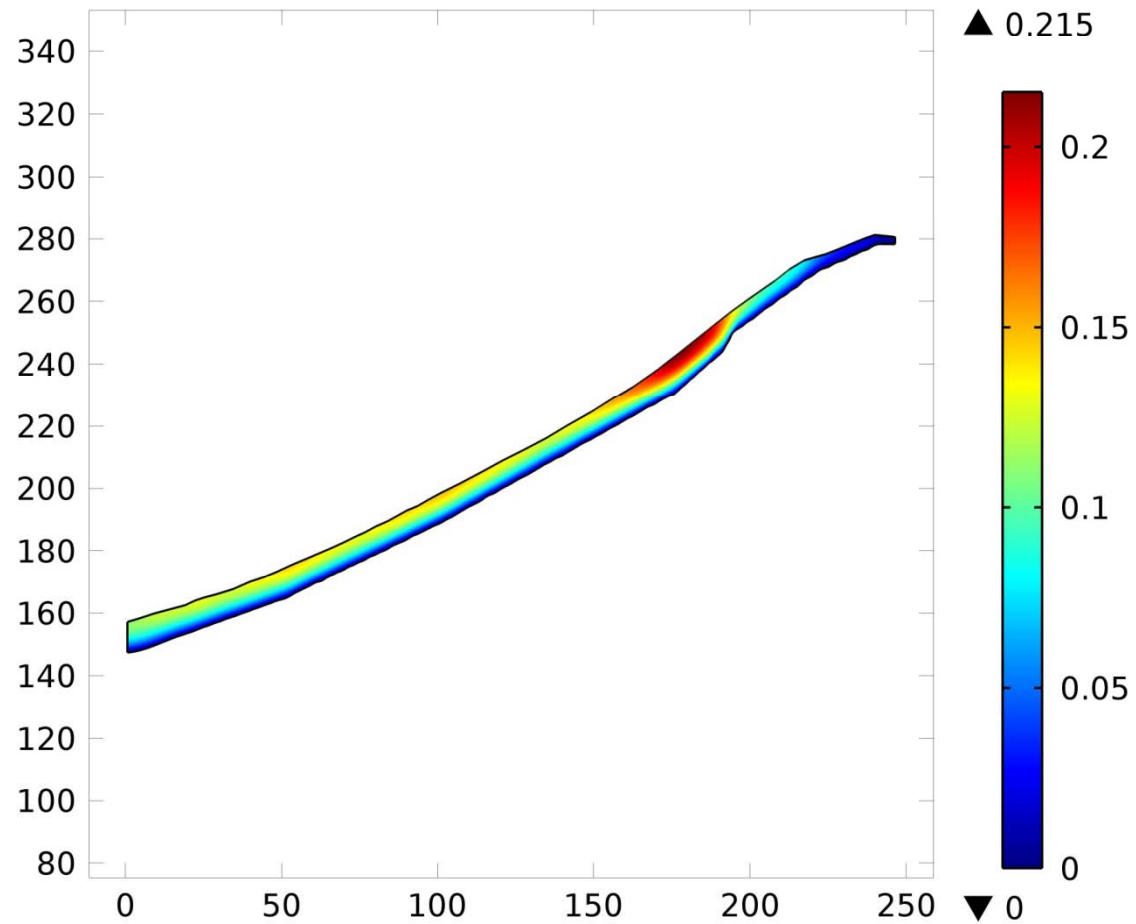
Degree of saturation on 9/28



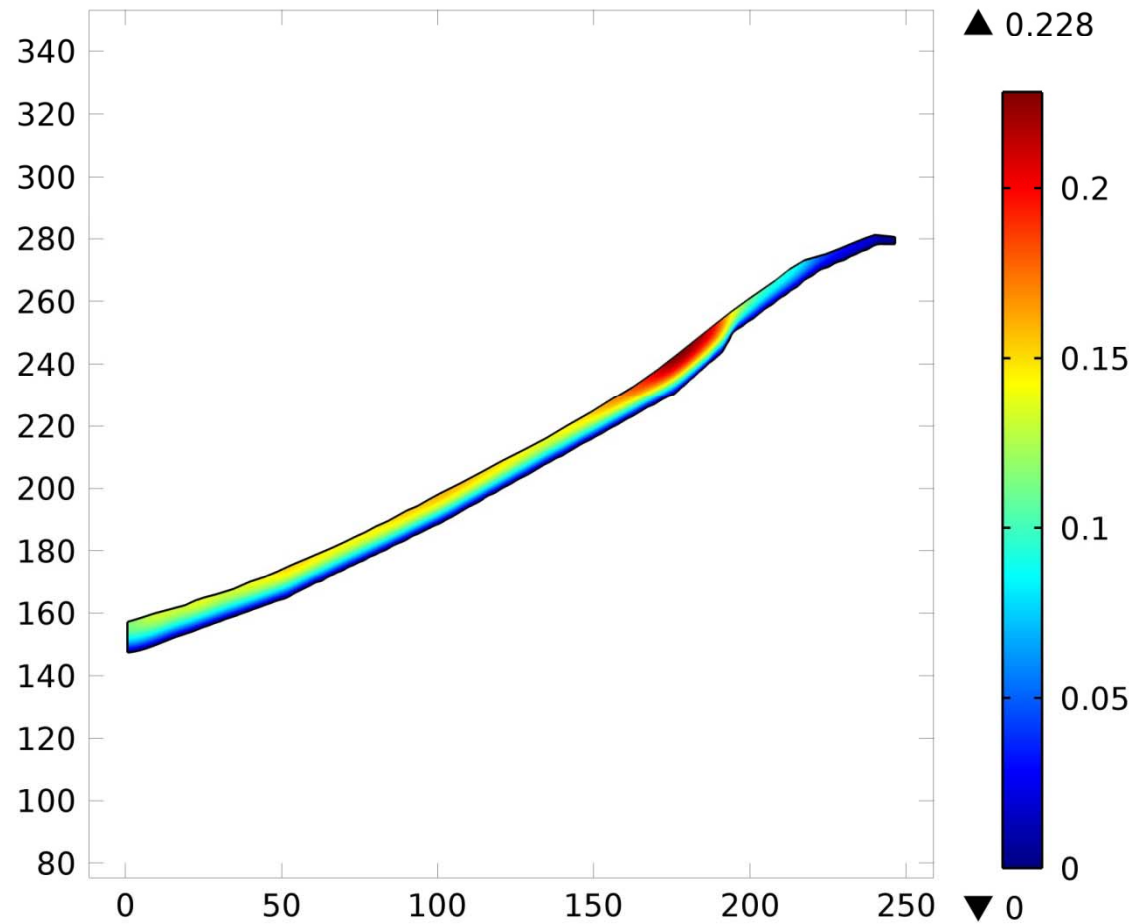
Displacement on 8/30



Displacement on 9/14



Displacement on 9/28



CONCLUSIONS

- A series of non linear PDEs has been used to asses the initiation of the mechanism of the Maokong landslide.
- Geotechincal problems related to contamination, groundwater flow, climate change, etc. can all be represented by non-linear PDEs and be solved using the increasingly powerful COMSOL.

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THE END